

# **Particle Tracking and the Diffusion-Reaction Equation**

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**Abstract.** Particle tracking algorithms are very useful methods to model conservative transport in surface and subsurface hydrological systems. Recently, a novel *ad hoc* particle-based method was proposed to account for multi-component reactive transport by *Benson and Meerschaert* [2008]. This one-dimensional particle method has been shown to match theoretical predictions, but to date there has been no rigorous demonstration that the particle method actually matches the governing equations for chemical transport. We generalize this particle method to two-dimensional and three-dimensional systems and rigorously demonstrate that this particle method converges to the diffusion-reaction equation at the limit of infinitely small time step. We also investigate the numerical error associated with the method.

## 1. Introduction

Water that flows through aquifer and surface water bodies typically changes in chemical composition along a flow path. One of the mechanisms for this is chemical reaction, either with neighboring solids and biota, or with other dissolved constituents. An accurate prediction of the chemical evolution of a system requires models (for example, governing equations) that correctly describe the chemistry and physics of the reactions across many spatial and temporal scales. Examples of these systems include radionuclide transport [Ma et al., 2010; Greskowiak et al., 2010; Hammond and Lichtner, 2010], CO<sub>2</sub> sequestration [Strazisar et al., 2006; Audigane et al., 2007; Han et al., 2010], ore body generation [Raffenspreger, 1997; Schardt et al., 2001], hydrothermal systems [Lichtner, 1985; Steefel and Lasaga, 1994], chemical and biochemical remediation of contaminated aquifers [Molz and Widdowson, 1988; Knutson et al., 2005, 2007; Steefel et al., 2005; Mayer et al., 2001, 2002, 2006], and basin-scale diagenesis, reservoir rock and petroleum generation [Lee, 1997; Morse et al., 1997; Wilson et al., June, 2000; Jones and Xiao, 2006] to name a few, but the problem is not restricted to aqueous environments [Searle et al., 1998a, b; Monson and Kopelman, 2000]. In many natural systems, the rate of reactions between the chemical species is limited by their mixing, because the molecules can only react when they come into physical contact. In systems that are well-mixed, like a stirred laboratory beaker, the fluctuations in concentrations of the chemical species are negligible and the rate of reaction is spatially uniform. However, if the mixing is slow in comparison to the rate of the reaction, a mixing-limited regime will emerge. In this regime, the fluctuations in concentrations possess a significant magnitude, and gradually we may begin to observe

43 areas of the domain where one reacting species is depleted relative to the others. From  
44 a mathematical point of view, the fluctuations of concentrations become anticorrelated  
45 [*Bolster et al.*, 2012; *Tartakovsky et al.*, 2012].

46 This segregation of the system into such “islands”, in which the overall reaction rate  
47 is dictated by the reduced area of contact between reactants, was described from theo-  
48 retical and numerical viewpoints several decades ago [*Ovchinnikov and Zeldovich*, 1978;  
49 *Toussaint and Wilczek*, 1983; *Kang and Redner*, 1984]. These authors showed that, in a  
50 simple chemical system that was dominated by diffusive transport, the segregation into  
51 islands was self-generated. The depletion of one of the reacting species inside the islands  
52 largely limits the reactions to island interfaces. As a result, the overall reaction rate  
53 is slower than the one anticipated for a well-mixed system. A small-scale experiment  
54 confirmed these results [*Monson and Kopelman*, 2004]. The slow-down of reaction rate  
55 has practical importance, especially in the groundwater remediation context, where the  
56 choice of remediation strategy is based on accurate modeling of the anticipated decay rate  
57 of the contaminants. To this end, a Lagrangian technique was proposed that could be im-  
58 plemented at any scale [*Benson and Meerschaert*, 2008]. The algorithm accounts for the  
59 increased probability that nearby particles will have for co-location (mixing) and reaction.  
60 It is the exact calculation of reaction probability that allows explicit simulation of poor  
61 mixing and reactant segregation. These numerical results were shown to correspond to  
62 perturbed continuum reaction-diffusion equations [*Bolster et al.*, 2012; *Tartakovsky et al.*,  
63 2012]. *However, to date the Lagrangian method has never been shown to converge to a*  
64 *specific governing equation.* It is therefore a phenomenological procedure: the particle  
65 methods can be used to solve for diffusion-limited reaction, but exactly what is being

66 solved? And what do changes in the numerical parameters represent in terms of physical  
 67 properties and numerical error? Our goal here is to show that the Lagrangian method  
 68 proposed by *Benson and Meerschaert* [2008] does correspond to a continuum equation of  
 69 transport and reaction, and that the numerical errors can be quantified and controlled by  
 70 the user.

71 The classical (continuum) approach to describe the fate and transport of a chemical  
 72 species is the advection-diffusion-reaction equation. In the special case of constant advec-  
 73 tion, the system can be transformed by a Galilean transformation and described by the  
 74 simpler diffusion-reaction equation (DRE)

$$\partial C_i / \partial t = \nabla \cdot (D \nabla C_i) - r_i \quad (1)$$

75 where  $C_i$  is the concentration of the chemical species  $i$  [mol/L<sup>*d*</sup>],  $D$  is the diffusion co-  
 76 efficient [L<sup>2</sup>/T], and  $r_i$  is the net rate of decay of the species [mol/L<sup>*d*</sup>T]. Here,  $d$  is **the**  
 77 dimension of the system ( $d = 1, 2, 3$ ).

78 For the case of radioactive decay,  $A \rightarrow C$ , the rate is typically described by the linear  
 79 rate  $r_i = k_d C_A$ , and the equation can be solved by means of Lagrangian particle tracking in  
 80 a straightforward manner [*Sherman and Peskin*, 1986; *Kinzelbach*, 1987]. In the context of  
 81 this paper we are interested in more complicated case of a bi-molecular reaction described  
 82 by



83 For this type of reaction, the rate is usually modeled by the (nonlinear) law of mass  
 84 action, i.e.

$$r_i = k C_A C_B \quad (3)$$

85 where  $i = A, B$ , and  $k$  is the constant rate coefficient [L<sup>*d*</sup>/mol T].

The modeling of coupled advection-diffusion-reaction equations with bi-molecular reac-  
 tion rate (3) is the subject of extensive research [Dentz *et al.*, 2011]. Benson and Meer-  
 schaert [2008] described a novel particle tracking (PT) approach for bi-molecular reactive  
 flow in a one-dimensional case with zero advection and a constant diffusion coefficient, and  
 demonstrated that the PT approach is capable of capturing the continuum of well-mixed  
 to incomplete mixing regimes. In short, within each time step, the PT approach moves  
 particles by random walk and then annihilates part of them at a probability related to  
 their distance apart. A novel feature of the PT approach is that it is purely Lagrangian  
 and grid-less. As such, it can accurately model incomplete mixing at all scales. This is  
 in contrast with approaches that assume complete mixing within the domain [Gillespie,  
 2000] or within volume elements [Isaacson, 2008]. In addition, the PT approach has an  
 advantage over the particle-grid approach of Thompson and Dougherty [1992], because the  
 latter involves calculation of the concentrations at grid points at every time step in order  
 to compute the reaction term. This is no longer needed within the purely PT scheme,  
 and the interpolation error involved with such a calculation is avoided. The analogy  
 between a random walk and the diffusion equation is well known [e.g., Einstein, 1906;  
 Chandrasekhar, 1943; Kinzelbach, 1987]. It remains to be shown that the addition of the  
 particle annihilation within the PT approach reflects the reaction term correctly. In other  
 words, we need to show that there is a mathematical analogy between the PT approach  
 and the DRE (1) in the limit of an infinitely small time step. Furthermore, we generalize  
 the PT approach to arbitrary dimension  $d$  and try to clarify some of the aspects of the  
 PT approach, with special care regarding numerical errors and the meaning of the initial  
 condition, namely the initial particle numbers.

## 2. Description of the PT simulation

The PT approach purports to solve (1) for a specified domain with appropriate initial and boundary conditions. In the PT approach, the particles of species  $A$  represent the concentration of that species,  $C_A$ , in a statistical manner. In other words, the specific locations of the particles in a specific simulation are usually meaningless; it is the ensemble average and other spatial and temporal statistical properties that are investigated.

### 2.1. Initialization of the PT

The primary advantage of the PT approach over deterministic approaches is in cases where the initial condition is non-deterministic and is given in terms of statistical parameters. A good example for this case is a system with an initial condition of a uniform mean concentration with some noise around it. Suppose that the mean concentration is equal for both species and given by

$$\langle C_i(x, t = 0) \rangle = C_0 \quad (4)$$

where  $\langle \cdots \rangle$  denotes the ensemble average,  $i = A, B$ , and  $C_0$  is a uniform prescribed initial concentration over the entire domain  $\Omega^d$ . Initially, the concentration fluctuations over the ensemble average, i.e.  $C'_i = C_i - \langle C_i \rangle$ , are assumed to have a short-range correlation that can be mathematically described by a Dirac-delta function,

$$\langle C'_i(\mathbf{x}, t = 0) C'_i(\mathbf{y}, t = 0) \rangle = \sigma^2 l^d \delta(\mathbf{x} - \mathbf{y}). \quad (5)$$

We note that the cross-correlation of the fluctuations is determined by the constant  $\sigma$ , denoting the magnitude of the concentration fluctuations [mol/L <sup>$d$</sup> ], and the constant  $l$ , denoting the length scale of these fluctuations [L]. The use of Dirac-delta function is known to be a good approximation of short-range correlations, such as Gaussian correlation and

127 exponential correlation [*Neuweiler et al.*, 2003; *Bolster et al.*, 2009; *Tartakovsky et al.*,  
128 2012].

129 For a domain of volume  $\Omega^d$ , this initial condition is modeled in the PT system by  
130 spreading  $N_0$  particles in the domain in a random manner. It has been shown by statistical  
131 analysis [*Paster et al.*] that the initial condition ((4)–(5)) dictates the initial density of  
132 the particles,

$$\frac{N_0}{\Omega^d} = \frac{C_0^2}{\sigma^2 l^d}. \quad (6)$$

133 In other words, the total number of particles is correlated to the initial condition: the  
134 more noise in the initial condition, the less particles need to be spread. Note that if the  
135 system is always completely smooth (i.e., perfectly mixed), then  $\sigma \rightarrow 0$ , and the problem  
136 may not be satisfactorily solved using the PT approach, because the large number of  
137 particles can be impractical. We define the mass of a single particle (the amount of moles  
138 described by a single numerical particle) by  $m_p$  [mol/particle]. As all particles carry the  
139 same mass, it is given by the total mass in the system divided by the number of particles,

$$m_p = C_0 \Omega^d / N_0. \quad (7)$$

140 Because we spread  $N_0$  particles for each species, we end up with a total number of  $2N_0$   
141 particles in the domain. Due to the random nature of the spread of the particles, there  
142 is no correlation initially between the locations of the  $A$  particles and the  $B$  particles,  
143 reflecting the initial condition

$$\langle C'_A(x, t = 0) C'_B(y, t = 0) \rangle = 0. \quad (8)$$



## 2.2. Description of PT steps

A single simulation time step in the PT approach describes the change of state of the system between  $t$  to  $t + \Delta t$ . It is composed of two sub-steps: random walk and reaction.

The random walk is an implementation of a Langevin equation,

$$x_j(t + \Delta t) = x_j(t) + \xi_j \sqrt{2D\Delta t} \quad (9)$$

where  $x_j$  is the  $j$ -th component of the particle location ( $j = 1, \dots, d$ ), and  $\xi_j$  is a random number of standard normal distribution, with zero mean and unit variance, i.e.  $\xi_j \sim \mathcal{N}(0, 1)$ . Thus, for each particle, its location is updated in all  $d$  dimensions by random values that account for the diffusion in this time step.

The reaction is then implemented by sequentially looping through all possible  $AB$  couples of particles in the system. Here, an  $AB$  couple is any combination of one  $A$  particle and one  $B$  particle (order being immaterial). For each such couple, a probability of forward reaction during the time step between those particles is then given by

$$p_f = km_p \Delta t v(s) \quad (10)$$

where  $s$  is the distance between the particles, and  $v(s)$  is the co-location probability density function (pdf). This is the pdf for two particles to co-locate over the time step, which for Brownian motion diffusion with coefficient  $D$  is given by [Benson and Meerschaert, 2008]:

$$v(s) = \frac{1}{(8\pi D\Delta t)^{d/2}} e^{-\frac{s^2}{8D\Delta t}}. \quad (11)$$

The probability of reaction, calculated by (10), is the probability of the  $AB$  couple to annihilate in this time step. Technically, this is done by producing a random number of uniform distribution  $\xi \sim U(0, 1)$ , and comparing these probabilities. Then, if  $p_f > \xi$ , the

particles are annihilated, i.e. removed from the system. The choice of time step must be such that  $p_f < 1$  is assured.

If one is interested in the fate of the product of the reaction (or in the case of backward reaction), a single  $C$  particle is placed into the domain. The location of this particle is randomly distributed, with the mean location at the mid point between the annihilated  $A, B$  particles, and with a variance of  $D\Delta t$ . It may be advected and diffused by a Langevin equation just as the  $A, B$  particles, and, depending on the reversibility of reaction, have a certain probability to decay back into  $A, B$  particles in some future time step.

The PT steps are repeated until the time for simulation end is approached, or, in the case of irreversible reaction, until all particles are consumed by the reaction. The PT simulation is repeated in a Monte Carlo fashion, so that results can be statistically analyzed with any specified degree of confidence. Typically, one would be interested in determining the average concentration in the system as a function of time, but other parameters, such as the concentration variance, or higher moments, can be computed as well.

### 3. The analogy of the PT approach with the DRE

We now show the analogy between the numerical PT scheme and the DRE equation. Start by defining  $W_A(\mathbf{x}, t)d\mathbf{x}$  as the expected number of  $A$  particles in the infinitesimal volume  $d\mathbf{x}$  centered at  $\mathbf{x}$ . The expected density of particles is then given by  $W_A(\mathbf{x}, t)$ , such that  $\int_{\Omega^d} W_A(\mathbf{x}, t)d\mathbf{x} = N(t)$ , where  $N$  is the total number of the  $A$  particles. For  $t = 0$ ,  $N(t) = N_0$ .

An initial distribution  $W_A(x, t = 0)$  in a specific simulation is the sum of  $N_0$  delta functions, positioned randomly in the domain  $\Omega^d$ . At later times, the distribution is

smeared by the diffusion and reduced by the reactions. It is worth noting that the density distribution is similar in essence to a probability density function, in the sense that it is non-negative, but different in the sense that the integration over the distribution does not sum to unity, but rather to  $N(t)$ .

After the particles have moved by the random walk, the expected particle density is the sum of the densities of all particles multiplied by the probability that they jump to location  $\mathbf{x}$ . This sum, neglecting boundary effects, is given by

$$W_A^*(\mathbf{x}, t + \Delta t) = \int W_A(\mathbf{y}, t) \frac{1}{(4\pi D \Delta t)^{d/2}} e^{-\frac{(\mathbf{x}-\mathbf{y})^2}{4D\Delta t}} d\mathbf{y} \quad (12)$$

where the star denotes the resulting intermediate distribution, before the reaction sub-step takes place. A similar expression is found for  $W_B^*$ , by replacing  $A$  by  $B$  in (12).

Next, the reaction takes place and is expected to annihilate a fraction of this particle density, such that the new distribution is given by

$$W_A(\mathbf{x}, t + \Delta t) = W_A^*(\mathbf{x}, t + \Delta t) - \Delta W_A^*(\mathbf{x}, t + \Delta t) \quad (13)$$

where  $\Delta W_A^*(\mathbf{x}, t + \Delta t)$  is the annihilated density. To express this density, we consider  $\Delta W_A^*(\mathbf{x}, t + \Delta t) \Delta \mathbf{x}$ , the number of particles in the volume  $[\mathbf{x}, \mathbf{x} + \Delta \mathbf{x})$  that will be annihilated. Clearly, if one  $A$  particle existed in this segment, and only one  $B$  particle would have been in its proximity, this probability would have been simply  $p_f$ , the forward reaction probability between two particles (10). However, neither the number of  $A$  particles in this segment, nor the number of nearby  $B$  particles, is necessarily unity. Therefore we need to multiply  $p_f$  by the actual number of  $A$  and  $B$  particles, and sum over the whole domain by means of integration. We can now write

$$\Delta W_A^*(\mathbf{x}, t + \Delta t) = W_A^*(\mathbf{x}, t + \Delta t) \int p_f(\mathbf{x} - \mathbf{x}') W_B^*(\mathbf{x}', t + \Delta t) d\mathbf{x}' \quad (14)$$

with  $p_f(\mathbf{x} - \mathbf{x}') = km_p v(\mathbf{x} - \mathbf{x}')\Delta t$ . Substituting (12) and (14) into (13), we get

$$\begin{aligned}
 W_A(\mathbf{x}, t + \Delta t) = & \\
 W_A^*(\mathbf{x}, t + \Delta t) & \left[ 1 - km_p \Delta t \int W_B^*(\mathbf{x}', t + \Delta t) v(\mathbf{x} - \mathbf{x}') d\mathbf{x}' \right] = \\
 & \int W_A(\mathbf{y}, t) \frac{1}{\sqrt{4\pi D \Delta t}} e^{-\frac{(\mathbf{x}-\mathbf{y})^2}{4D\Delta t}} d\mathbf{y} \\
 & \times \left[ 1 - km_p \Delta t \int v(\mathbf{x} - \mathbf{x}') d\mathbf{x}' \int W_B(\mathbf{y}', t) \frac{1}{\sqrt{4\pi D \Delta t}} e^{-\frac{(\mathbf{x}'-\mathbf{y}')^2}{4D\Delta t}} d\mathbf{y}' \right] \quad (15)
 \end{aligned}$$

Expanding in a Taylor series for  $W_A(\mathbf{x}, t)$  we get, for the one dimensional case

$$W_A(y, t) = W_A(x, t) + \Delta x \frac{\partial W_A}{\partial x} + \frac{(\Delta x)^2}{2!} \frac{\partial^2 W_A}{\partial x^2} + \dots \quad (16)$$

For the two-dimensional case, we get

$$\begin{aligned}
 W_A(\mathbf{y}, t) = W_A(\mathbf{x}, t) + \Delta x_1 \frac{\partial W_A}{\partial x_1} + \Delta x_2 \frac{\partial W_A}{\partial x_2} + \\
 \frac{1}{2!} \left[ (\Delta x_1)^2 \frac{\partial^2 W_A}{\partial x_1^2} + 2\Delta x_1 \Delta x_2 \frac{\partial^2 W_A}{\partial x_1 \partial x_2} + (\Delta x_2)^2 \frac{\partial^2 W_A}{\partial x_2^2} \right] + \dots \quad (17)
 \end{aligned}$$

where  $\Delta \mathbf{x} = \mathbf{y} - \mathbf{x} = (\Delta x_1, \Delta x_2)$ . For the three-dimensional case,

$$\begin{aligned}
 W_A(\mathbf{y}, t) = W_A(\mathbf{x}, t) + \Delta x_j \frac{\partial W_A}{\partial x_j} + \frac{1}{2!} \left[ (\Delta x_j)^2 \frac{\partial^2 W_A}{\partial x_j^2} + \right. \\
 \left. 2\Delta x_1 \Delta x_2 \frac{\partial^2 W_A}{\partial x_1 \partial x_2} + 2\Delta x_1 \Delta x_3 \frac{\partial^2 W_A}{\partial x_1 \partial x_3} + 2\Delta x_2 \Delta x_3 \frac{\partial^2 W_A}{\partial x_2 \partial x_3} \right] + \dots \quad (18)
 \end{aligned}$$

where Einstein notation is applied and  $\Delta \mathbf{x} = (\Delta x_1, \Delta x_2, \Delta x_3) = \mathbf{y} - \mathbf{x}$ . Substitution of

the appropriate Taylor expansion (16-18) into the integral in (15) yields

$$\int W_A(\mathbf{y}, t) \frac{1}{(4\pi D \Delta t)^{d/2}} e^{-\frac{(\mathbf{x}-\mathbf{y})^2}{4D\Delta t}} d\mathbf{y} = W_A(\mathbf{x}, t) + D\Delta t \frac{\partial^2 W_A}{\partial x_j \partial x_j} + \mathcal{O}((D\Delta t)^2) \quad (19)$$

and similarly

$$\int W_B(\mathbf{y}', t) \frac{1}{(4\pi D \Delta t)^{d/2}} e^{-\frac{(\mathbf{x}'-\mathbf{y}')^2}{4D\Delta t}} d\mathbf{y}' = W_B(\mathbf{x}', t) + D\Delta t \frac{\partial^2 W_B}{\partial x_j \partial x_j} + \mathcal{O}((D\Delta t)^2). \quad (20)$$

By substituting (19) and (20) into (15) we get

$$\begin{aligned}
 W_A(\mathbf{x}, t + \Delta t) = & \left[ W_A(\mathbf{x}, t) + D\Delta t \frac{\partial^2 W_A}{\partial x_j \partial x_j} + \dots \right] \times \\
 & \left\{ 1 - km_p \Delta t \int v(\mathbf{x} - \mathbf{x}') d\mathbf{x}' \left[ W_B(\mathbf{x}', t) + D\Delta t \frac{\partial^2 W_B}{\partial x_j \partial x_j} + \dots \right] \right\} = \\
 & W_A(\mathbf{x}, t) + D\Delta t \frac{\partial^2 W_A}{\partial x_j \partial x_j} - km_p \Delta t W_A(\mathbf{x}, t) \int v(\mathbf{x} - \mathbf{x}') W_B(\mathbf{x}', t) d\mathbf{x}' + \mathcal{O}((\Delta t)^2) \quad (21)
 \end{aligned}$$

Taking the integral on the RHS of (21), expanding  $W_B$  once more in a Taylor series as

in (16-18), and using (11), we find

$$\begin{aligned}
 \int v(\mathbf{x} - \mathbf{x}') W_B(\mathbf{x}', t) d\mathbf{x}' = \\
 \int \frac{1}{(8\pi D\Delta t)^{d/2}} e^{-\frac{(\mathbf{x}-\mathbf{x}')^2}{8D\Delta t}} W_B(\mathbf{x}', t) d\mathbf{x}' = \\
 W_B(\mathbf{x}, t) + 2D\Delta t \frac{\partial^2 W_B}{\partial x_j \partial x_j} + \dots \quad (22)
 \end{aligned}$$

Substituting (22) into (21), subtracting  $W_A(\mathbf{x}, t)$  from both sides, and dividing by  $\Delta t$ ,

we find

$$\frac{W_A(\mathbf{x}, t + \Delta t) - W_A(\mathbf{x}, t)}{\Delta t} = D \frac{\partial^2 W_A}{\partial x_j \partial x_j} - km_p W_A(\mathbf{x}, t) W_B(\mathbf{x}, t) + \mathcal{O}(\Delta t) \quad (23)$$

As  $\Delta t \rightarrow 0$ , this equation becomes

$$\frac{\partial W_A}{\partial t} = D \frac{\partial^2 W_A}{\partial x_j \partial x_j} - km_p W_A W_B. \quad (24)$$

Finally, we recognize that the density of the particles  $W_A$  (of units  $[L^{-d}]$ ) can be ex-

pressed in terms of concentration of the species (of units  $[mol/L^d]$ ) by simply multiplying

it by the mass of a single particle  $m_p$ , i.e.

$$C_A = W_A m_p. \quad (25)$$

Hence, multiplying both sides of (24) by  $m_p$  we get the one-dimensional DRE for constant  $D$ ,

$$\frac{\partial C_A}{\partial t} = D\nabla^2 C_A - kC_A C_B. \quad (26)$$

Thus, we have shown that the PT approach, at the limit of  $\Delta t \rightarrow 0$ , converges to the DRE.

#### 4. Discussion and Conclusions

The utility of the PT approach for non-reactive particles is clear: the approach does not suffer from numerical dispersion and is potentially highly parallelizable. The problem of using the approach for reactive particles was first addressed by *Tompson and Dougherty* [1992], who discretize space into bins, count particles, convert to concentrations, calculate classical reactions, and re-convert concentrations to particles. The PT method of *Benson and Meerschaert* [2008] obviates the particle/concentration conversion at each time step, and eliminates the assumption of perfect mixing at the bin (or Eulerian block) scale. On the other hand, the PT approach had not been linked definitively to a differential equation of transport and reaction. In this work we were able to show the analogy between the diffusion-reaction equation (DRE) and the PT approach. This was done at the limit of  $\Delta t \rightarrow 0$ , i.e. an infinitely small time step. To this aim, we defined a density function of the particles, and described its evolution over time due to random walks and reaction between particles. Using Taylor series expansions and basic mathematical concepts we were able to show that the density is described by the DRE at the  $\Delta t \rightarrow 0$  limit. This was done for an arbitrary physical dimension ( $d = 1, 2, 3$ ). We also showed that the error induced by the scheme is  $\mathcal{O}(\Delta t)$ .

This analysis strictly speaking is limited to zero (or spatially constant) advection, whereas transport in natural porous media is typically characterized by a velocity field that can vary in space. Derivation of a particle tracking approach for solving the advection-diffusion-reaction in porous media is therefore a very important next step that is beyond the scope of this paper. It appears that the appropriate approach for this problem is to perform operator splitting, i.e., take the advective, diffusive, and reactive steps sequentially. As shown theoretically and experimentally by *Taylor* [1953], advection, even when heterogeneous, does not cause mixing of solutes, only diffusion does. Therefore, as long as the time step is sufficiently small that any errors induced by operator splitting are small, it should be sufficient to show that the reaction-diffusion part of the particle approach is correctly modeling the equation. This remains to be proven.

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## References

Audigane, P., I. Gaus, I. Czernichowski-Lauriol, K. Pruess, and T. Xu (2007), Two-dimensional reactive transport modeling of CO<sub>2</sub> injection in a saline aquifer at the Sleipner site, North Sea, *Am. J. Sci.*, *307*, 974–1008.

- 257 Benson, D. A., and M. M. Meerschaert (2008), Simulation of chemical reaction via parti-  
258 cle tracking: Diffusion-limited versus thermodynamic rate-limited regimes, *Water Re-*  
259 *sources Research*, *44*, W12,201.
- 260 Bolster, D., M. Dentz, and J. Carrera (2009), Effective two-phase flow in heterogeneous  
261 media under temporal pressure fluctuations, *Water Resour. Res.*, *45*, W05,408.
- 262 Bolster, D., P. de Anna, D. A. Benson, and A. M. Tartakovsky (2012), Incomplete mixing  
263 and reactions with fractional dispersion, *Advances in Water Resources*, *37*, 86 – 93.
- 264 Chandrasekhar, S. (1943), Stochastic problems in physics and astronomy, *Rev. Mod. Phys.*,  
265 *15*, 1–89.
- 266 Dentz, M., T. Le Borgne, A. Englert, and B. Bijeljic (2011), Mixing, spreading and  
267 reaction in heterogeneous media: A brief review, *Journal of Contaminant Hydrology*,  
268 *120-21*, 1–17.
- 269 Einstein, A. (1906), On the theory of the Brownian movement, *Annalen der Physik*, *4*,  
270 371–381.
- 271 Gillespie, D. (2000), The chemical langevin equation, *Journal of Chemical Physics*, *113*(1),  
272 297–306.
- 273 Greskowiak, J., H. Prommer, C. Liu, V. E. A. Post, R. Ma, C. Zheng, and J. M.  
274 Zachara (2010), Comparison of parameter sensitivities between a laboratory and field-  
275 scale model of uranium transport in a dual domain, distributed rate reactive system,  
276 *Water Resour. Res.*, *46*, W09,509.
- 277 Hammond, G. E., and P. C. Lichtner (2010), Field-scale model for the natural attenuation  
278 of uranium at the Hanford 300 area using high-performance computing, *Water Resour.*  
279 *Res.*, *46*, W09,527.



- 280 Han, W. S., B. J. McPherson, P. C. Lichtner, and F. P. Wang (2010), Evaluation of  
281 trapping mechanisms in geologic CO<sub>2</sub> sequestration: Case study of SACROC northern  
282 platform, a 35-year CO<sub>2</sub> injection site, *Am. J. Sci.*, *310*, 282–324.
- 283 Isaacson, S. A. (2008), Relationship between the reaction-diffusion master equation and  
284 particle tracking models, *Journal of Physics A: Mathematical and Theoretical*, *41*,  
285 065,003.
- 286 Jones, G. D., and Y. Xiao (2006), Geothermal convection in the Tengiz carbonate plat-  
287 form, Kazakhstan: Reactive transport models of diagenesis and reservoir quality, *AAPG*  
288 *Bulletin*, *90*, 1251–1272.
- 289 Kang, K., and S. Redner (1984), Scaling approach for the kinetics of recombination pro-  
290 cesses, *Physical Review Letters*, *52*, 955.
- 291 Kinzelbach, W. (1987), The random walk method in pollutant transport simulation, in  
292 *Groundwater Flow and Quality Modelling, NATO ASI Series C: Mathematical and*  
293 *Physical Sciences*, vol. 224, edited by E. Custodio, A. Gurgui, and J. L. Ferreira, pp.  
294 227–245.
- 295 Knutson, C., A. Valocchi, and C. Werth (2007), Comparison of continuum and pore-  
296 scale models of nutrient biodegradation under transverse mixing conditions, *Advances*  
297 *in Water Resources*, *30*(67), 1421 – 1431, biological processes in porous media: From  
298 the pore scale to the field.
- 299 Knutson, C. E., C. J. Werth, and A. J. Valocchi (2005), Pore-scale simulation of biomass  
300 growth along the transverse mixing zone of a model two-dimensional porous medium,  
301 *Water Resour. Res.*, *41*(7), W07,007.

302 Lee, M.-K. (1997), Predicting diagenetic effects of groundwater flow in sedimentary basins:  
303 a modeling approach with examples, in *Basin-Wide Diagenetic Patterns*, vol. 57, pp.  
304 3–14, SEPM (Society for Sedimentary Geology).

305 Lichtner, P. C. (1985), Continuum model for simultaneous chemical reactions and mass  
306 transport in hydrothermal systems, *Geochimica et Cosmochimica Acta*, 49(3), 779–800.

307 Ma, R., C. Zheng, H. Prommer, J. Greskowiak, C. Liu, J. Zachara, and M. Rockhold  
308 (2010), A field-scale reactive transport model for U(VI) migration influenced by coupled  
309 multirate mass transfer and surface complexation reactions, *Water Resour. Res.*, 46,  
310 W05,509.

311 Mayer, K., S. Benner, E. Frind, S. Thornton, and D. Lerner (2001), Reactive transport  
312 modeling of processes controlling the distribution and natural attenuation of phenolic  
313 compounds in a deep sandstone aquifer, *Journal of Contaminant Hydrology*, 53(34),  
314 341–368.

315 Mayer, K., S. Benner, and D. Blowes (2006), Process-based reactive transport modeling of  
316 a permeable reactive barrier for the treatment of mine drainage, *Journal of Contaminant*  
317 *Hydrology*, 85(34), 195 – 211.

318 Mayer, K. U., E. O. Frind, and D. W. Blowes (2002), Multicomponent reactive trans-  
319 port modeling in variably saturated porous media using a generalized formulation for  
320 kinetically controlled reactions, *Water Resour. Res.*, 38, 1174.

321 Molz, F., and M. Widdowson (1988), Internal inconsistencies in dispersion-dominated  
322 models that incorporate chemical and microbial kinetics, *Water Resour. Res.*, 24, 615–  
323 619.

- 324 Monson, E., and R. Kopelman (2000), Observation of laser speckle effects and nonclassical  
325 kinetics in an elementary chemical reaction, *Physical Review E*, *85*, 666–669.
- 326 Monson, E., and R. Kopelman (2004), Nonclassical kinetics of an elementary  $a + b \rightarrow c$   
327 reaction-diffusion system showing effects of a speckled initial reactant distribution and  
328 eventual self-segregation: Experiments, *Physical Review E*, *69*, 021,103.
- 329 Morse, J. W., J. S. Hanor, and S. He (1997), The role of mixing and migration of basinal  
330 waters in carbonate mineral mass transport, in *Basin-Wide Diagenetic Patterns*, vol. 57,  
331 pp. 41–50, SEPM (Society for Sedimentary Geology).
- 332 Neuweiler, I., S. Attinger, W. Kinzelbach, and P. King (2003), Large scale mixing for  
333 immiscible displacement in heterogeneous porous media, *Transp. Porous Media*, *51*,  
334 287–314.
- 335 Ovchinnikov, A., and Y. Zeldovich (1978), Role of density fluctuations in bimolecular  
336 reaction kinetics, *Chemical Physics*, *28*, 215 – 218.
- 337 Paster, A., D. Bolster, and D. A. Benson (), Connecting the dots: application of a particle  
338 method to the diffusion-reaction equation, submitted to *Advances in Water Resources*.
- 339 Raffenspreger, J. P. (1997), Evidence and modeling of large-scale groundwater convection  
340 in precambrian sedimentary basins, in *Basin-Wide Diagenetic Patterns*, vol. 57, pp.  
341 15–26, SEPM (Society for Sedimentary Geology).
- 342 Schardt, C., D. R. Cooke, J. B. Gemmell, and R. R. Large (2001), Geochemical modeling  
343 of the zoned footwall alteration pipe, Hellyer volcanic-hosted massive sulfide deposit,  
344 western Tasmania, Australia, *Econ. Geol.*, *96*, 1037–1054.
- 345 Searle, K. R., M. P. Chipperfield, S. Bekki, and J. A. Pyle (1998a), The impact of spatial  
346 averaging on calculated polar ozone loss 1. model experiments, *J. Geophys. Res.*, *103*,

25,397–25,408.

Searle, K. R., M. P. Chipperfield, S. Bekki, and J. A. Pyle (1998b), The impact of spatial averaging on calculated polar ozone loss 2. theoretical analysis, *J. Geophys. Res.*, *103*(D19), 25,409–25,416.

Sherman, A. S., and C. S. Peskin (1986), A Monte Carlo method for scalar reaction diffusion equations, *SIAM J. Sci. Stat. Comput.*, *7*(4), 1360–1372.

Steefel, C. I., and A. C. Lasaga (1994), A coupled model for transport of multiple chemical species and kinetic precipitation/dissolution reactions with application to reactive flow in single phase hydrothermal systems, *Am. J. Sci.*, *294*(5), 529–592.

Steefel, C. I., D. J. DePaolo, and P. C. Lichtner (2005), Reactive transport modeling: An essential tool and a new research approach for the earth sciences, *Earth and Planetary Science Letters*, *240*(34), 539–558.

Strazisar, B. R., C. Zhu, and S. W. Hedges (2006), Preliminary modeling of the long-term fate of CO<sub>2</sub> following injection into deep geological formations, *Environ. Geosci.*, *13*(1), 1–15.

Tartakovsky, A. M., P. de Anna, T. L. Borgne, A. Balter, and D. Bolster (2012), Effect of spatial concentration fluctuations on effective kinetics in diffusion-reaction systems, *Water Resources Research*, *48*, W02,526.

Taylor, G. (1953), Dispersion of soluble matter in solvent flowing slowly through a tube, *Proceedings of the Royal Society of London. Series A. Mathematical and Physical Sciences*, *219*(1137), 186–203, doi:10.1098/rspa.1953.0139.

Tompson, A., and D. Dougherty (1992), Particle-grid methods for reacting flows in porous-media with application to Fisher equation, *Appl. Math. Model.*, *16*(7), 374–383.

- 370 Toussaint, D., and F. Wilczek (1983), Particle-antiparticle annihilation in diffusive motion,  
371 *J. Chem. Phys.*, 78, 2642–2647.
- 372 Wilson, A. M., J. R. Boles, and G. Garven (June, 2000), Calcium mass transport and  
373 sandstone diagenesis during compaction-driven flow: Stevens Sandstone, San Joaquin  
374 basin, California, *Geol. Soc. Amer. Bull.*, 112, 845–856.